
Electrical Surveying

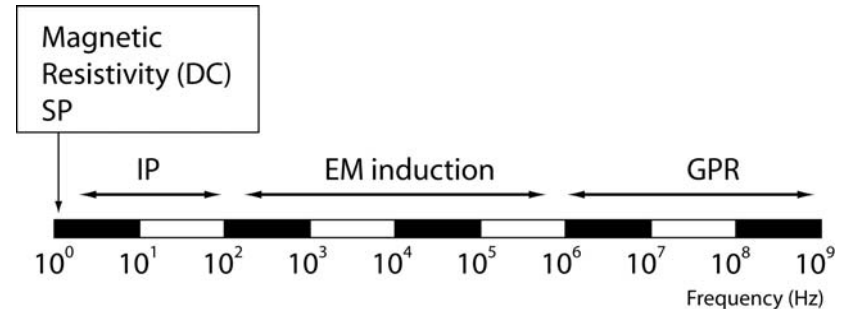
Part III: Self-potential method

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Introduction

Electrical surveying...



- Resistivity method
- Induced polarization method (IP)
- Self-potential (SP) method

Higher frequency methods (electromagnetic surveys):

- Electromagnetic induction methods
- Ground penetrating radar (GPR)

Self-potential method

The self-potential method makes use of **natural currents** flowing in the ground that are generated by **electrochemical processes** to locate shallow bodies of anomalous conductivity and water circulation

Application

- Exploration of metalliferous mineral deposits
- Detection of water circulation into the ground
- Monitoring dams or tank integrity

Structure of the lecture

1. Basic SP theory
2. Survey strategies and interpretation
3. Conclusions



1. Basic SP theory

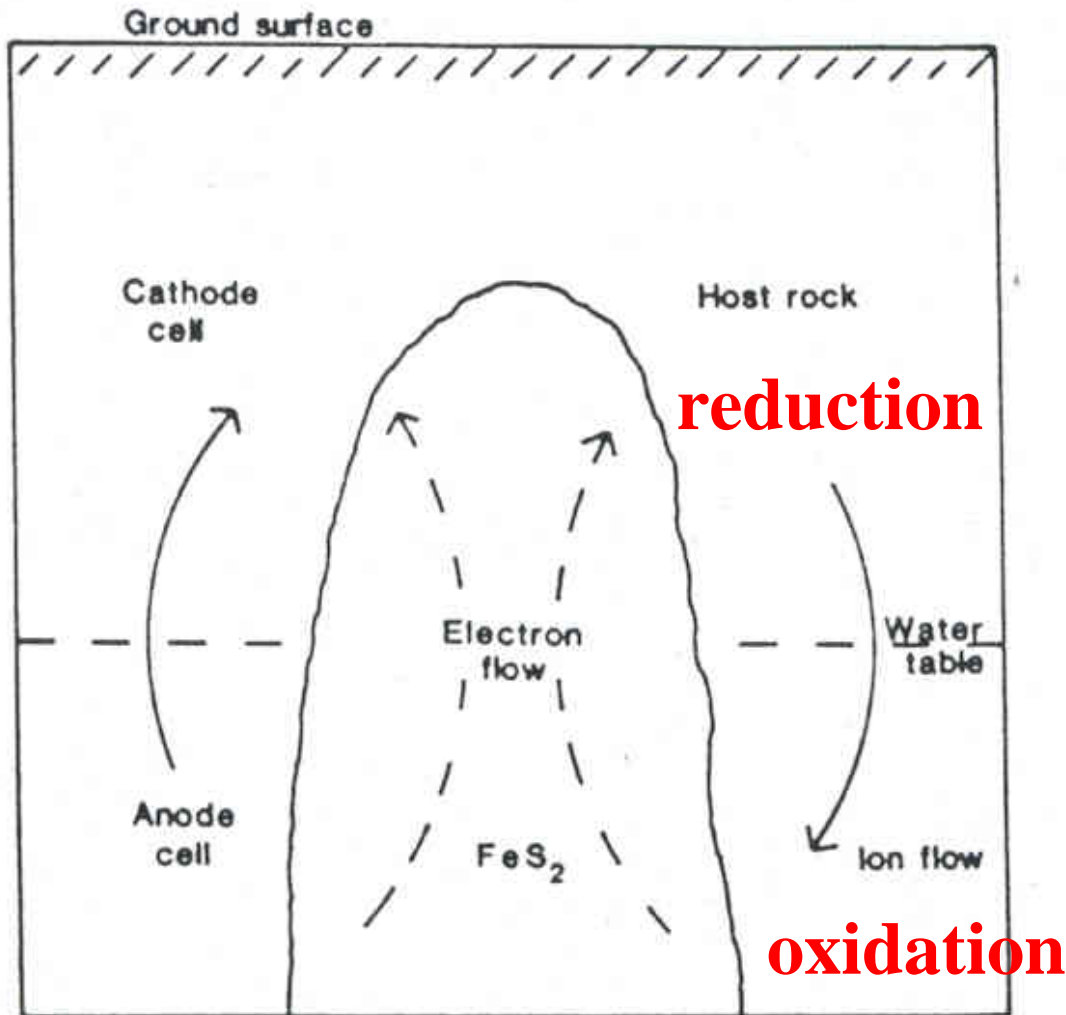
Basic theory

This is a method employing natural electric sources. Two natural potentials are mainly used in exploration geophysics:

- Electrokinetic (streaming) potential
- Mineralization potential

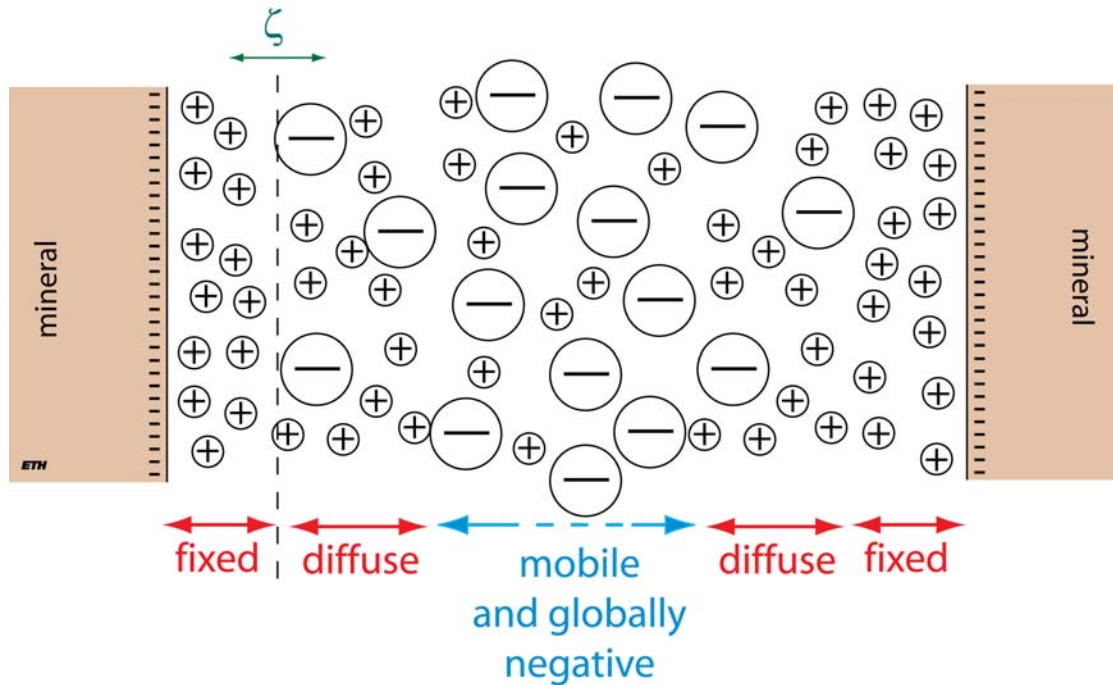
These potentials are expressed in V or mV

Mineralization potential



- Mining geophysics
- Negative anomalies often higher than 100 mV

Electrokinetic (streaming) potential



$$\Delta V = \frac{\zeta k}{\eta \sigma_w} \Delta P$$

ζ potential between + and - layers

k dielectric constant of the fluid

η viscosity of the fluid

σ_w conductivity of the fluid

ΔP difference of pressure between the measuring points

ΔV difference of potential between the ends of the passage

- Mainly used in hydrogeophysics
- Positive or negative anomalies often lower than 50 mV

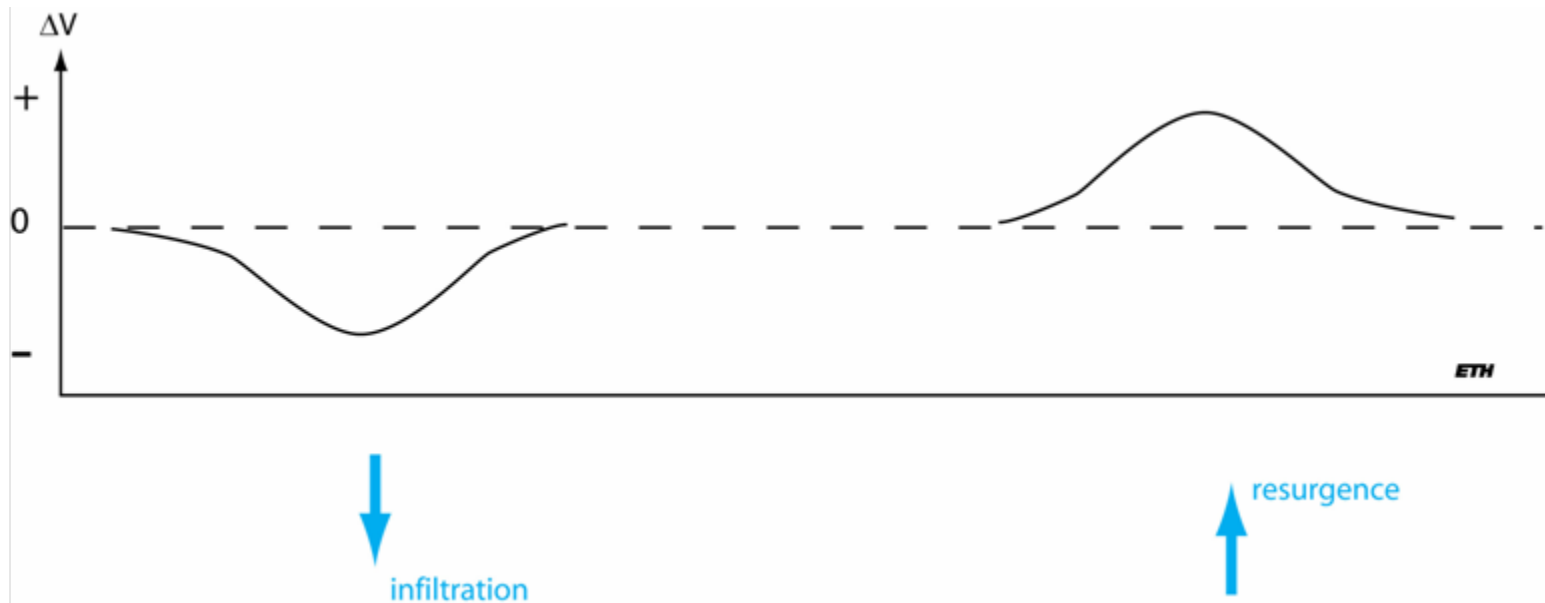
Electrokinetic (streaming) potential

$$C = \frac{\Delta V}{\Delta P} = \frac{\zeta k}{\eta \sigma_w}$$

C is often described as the coupling coefficient

- If the grain size decreases, C increases
- If the temperature decreases, C decreases
- If σ_w decreases, C increases
- Permeability has a complex effect on C .

Electrokinetic (streaming) potential



SP anomaly is positive where the hydrostatic pressure decreases (i.e. in the direction of the water flux).

If the flow is horizontal, then $\Delta P = \Delta h = 0$ and $\Delta V = 0$

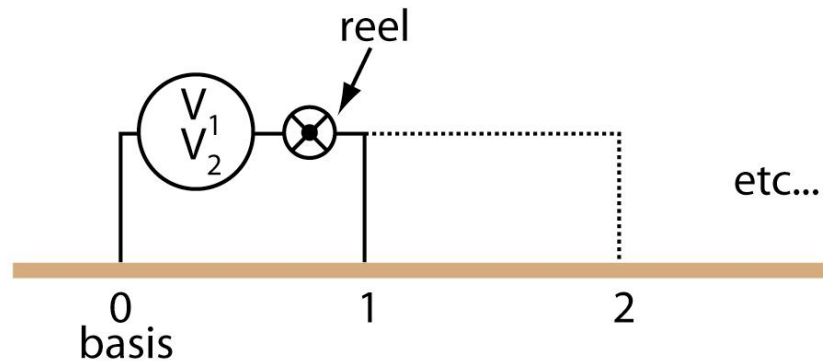
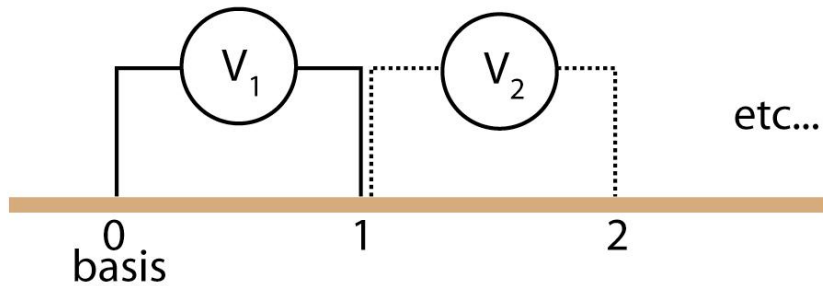


2. Survey strategies and interpretation

SP measurement

V : raw data

P : real potential values at the basis



$$P_1 = V_1$$

$$P_2 = V_1 + V_2$$

$$P_n = V_1 + V_2 + \dots + V_n$$

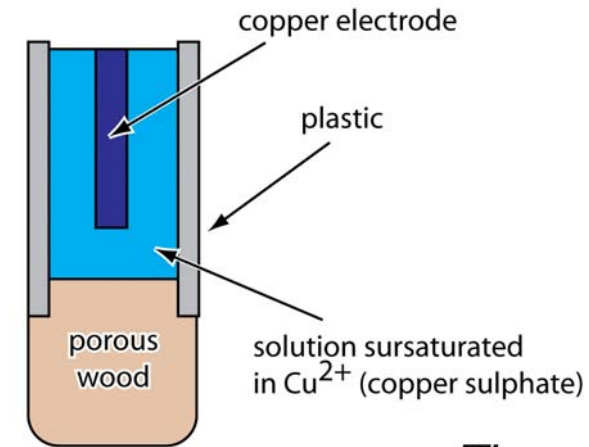
$$P_1 = V_1$$

$$P_2 = V_2$$

$$P_n = V_n$$

$$P_{n+1} = V_{n+1} + P_n \quad (\text{for a new basis})$$

Non-polarizing electrode

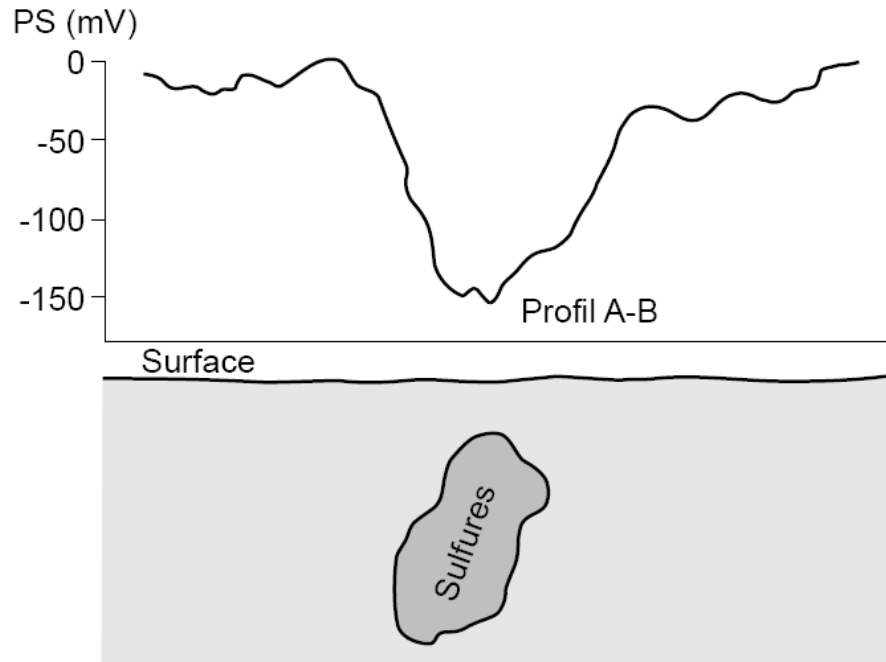


ETH

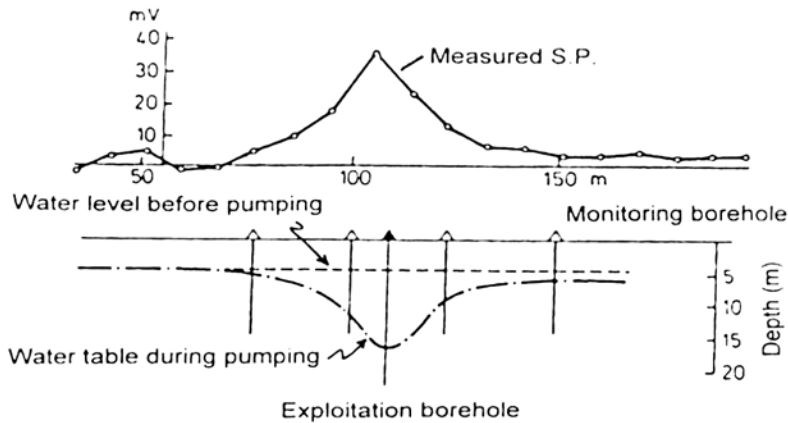
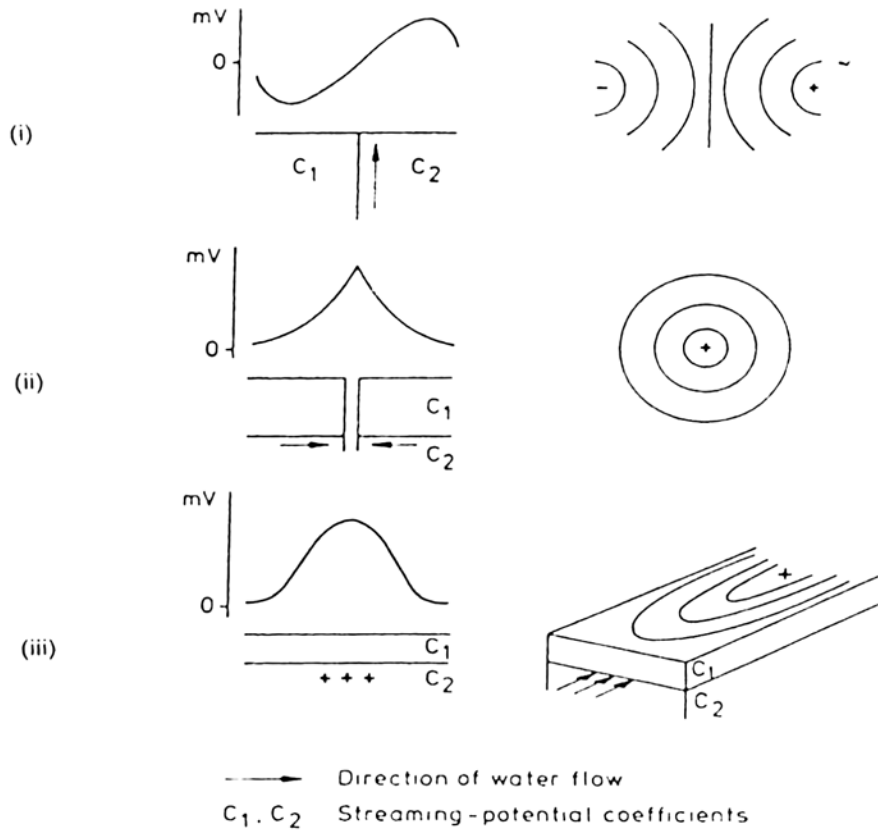
Interpretation

- Depth of investigation depends on the size of the mineralized body and the depth of the water table for a mineralization potential (generally shallow, < 30 m)
- Interpretation mainly qualitative (profile, map)
- Quantitative using dipole approximations for the polarized body (similar to magnetic interpretation)

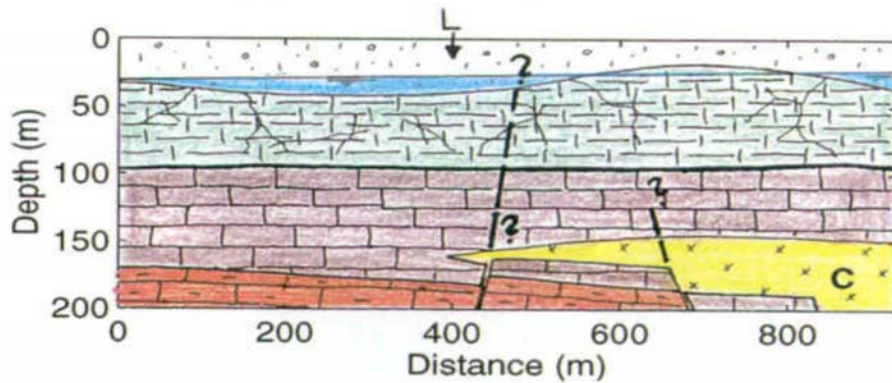
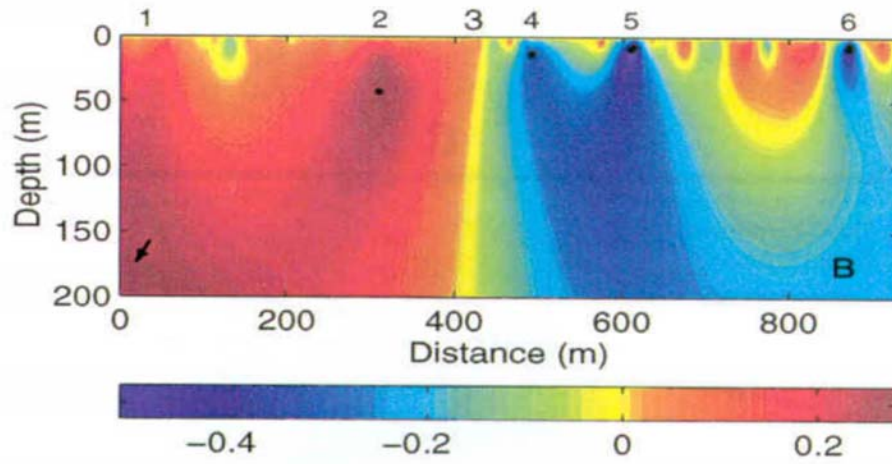
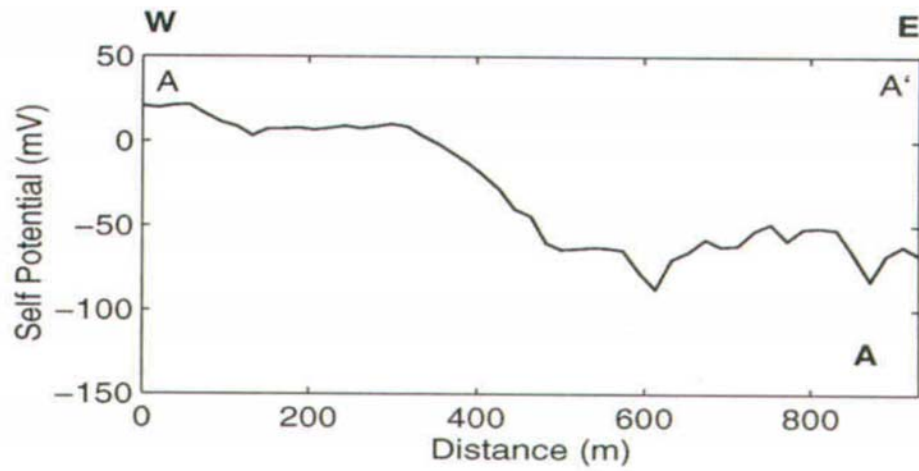
Example of mineralization potential



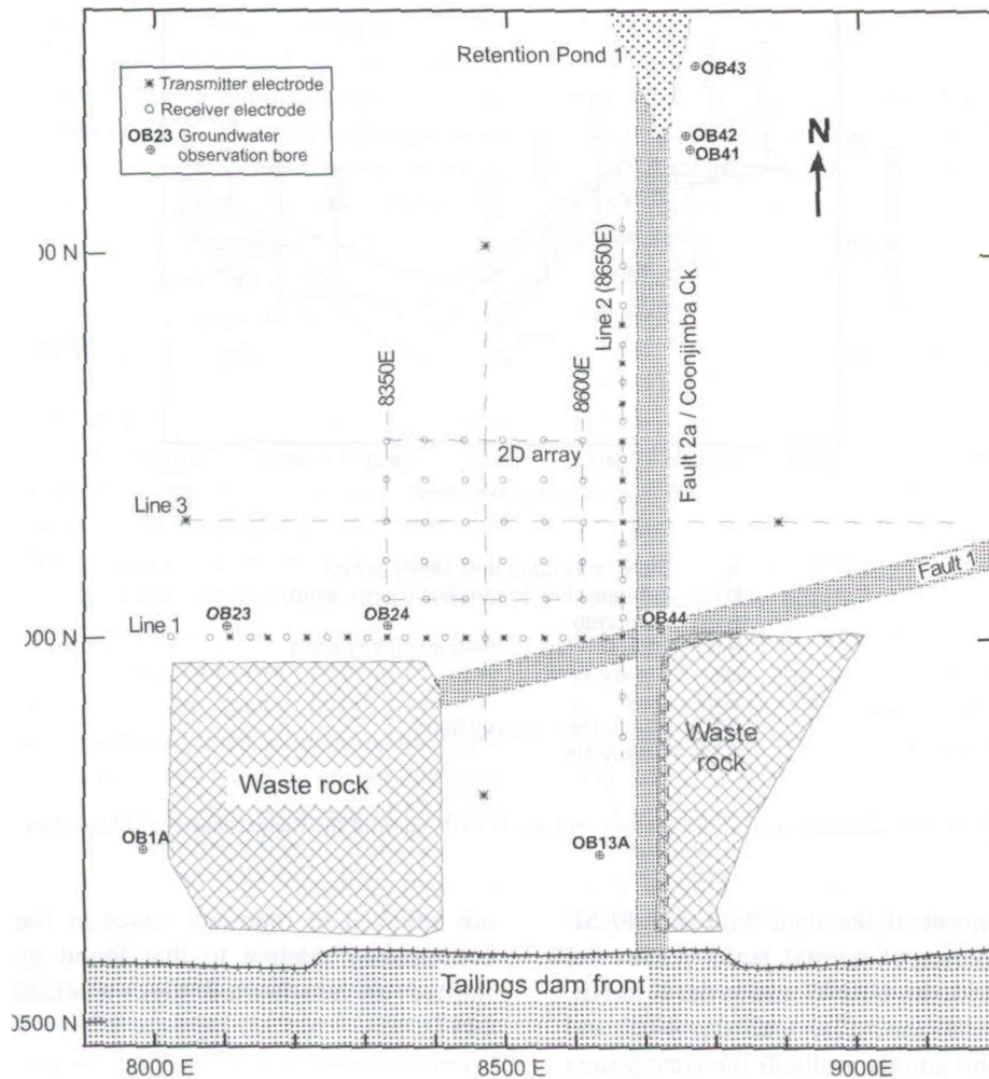
Examples of streaming potential



Geology



Waste deposits

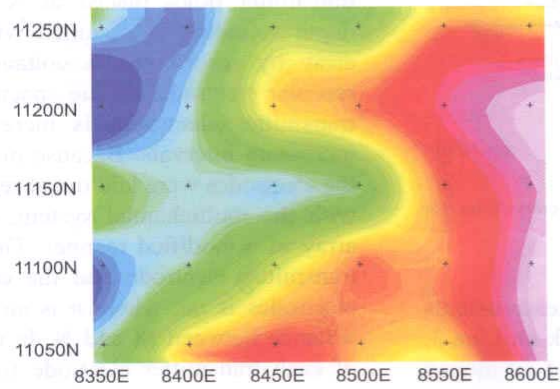


(a) The survey grid

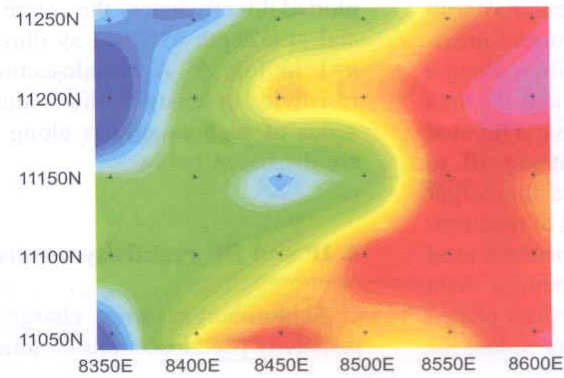


(b) Details of Line 3

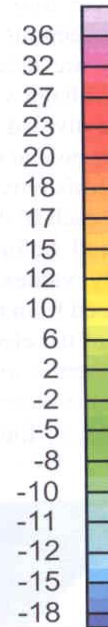
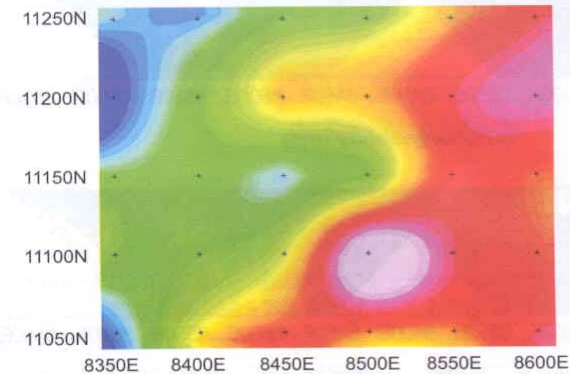
12:23pm, 23/10/98 (average of 10 sec data)



1:32pm, 23/10/98 (average of 200 sec data)

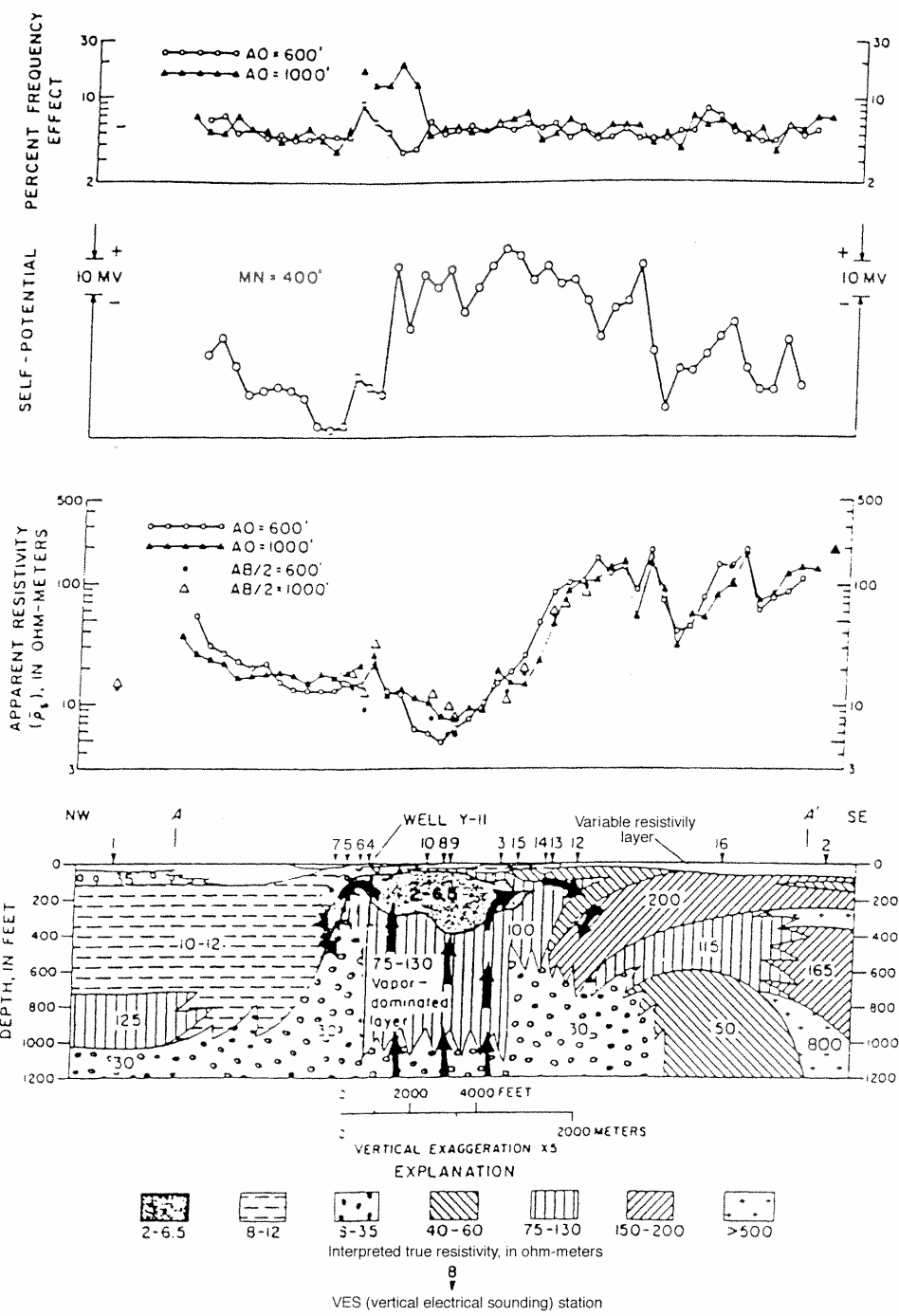


1:32pm, 26/10/98 (average of 600 sec data)



mV

Geothermy





3. Conclusions

Advantages

- Survey simple
- Non expensive
- Allows for a rapid qualitative mapping of the underground
- Suitable for monitoring

Drawbacks

- Very sensitive to noise
- Physical aspects still not well understood
- Quantitative aspects still need to be developed